







TECTÓNICA ACTIVA Y PALEOSISMOLOGÍA EN LA FALLA TAMBOMACHAY: en busca de los sismos

Sencca perdidos en el Cusco.

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Convenio 006-2016-Fondecyt: Proyecto Cusco-PATA

4180 msnm



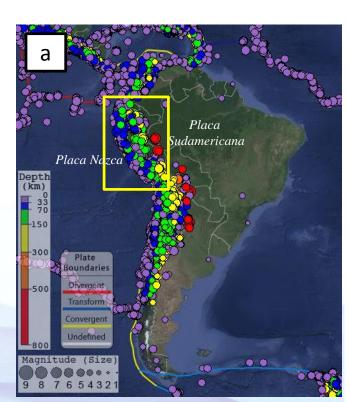


CONTENIDO

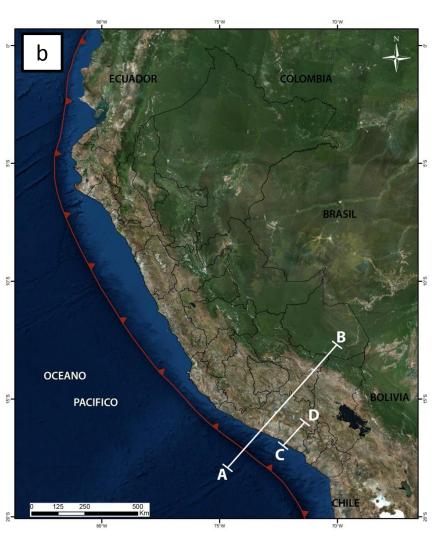
- ✓ Introducción
- ✓ Contexto Sismotectónico y geológico
- ✓ Estudio geomorfológico (metodología y resultados)
- ✓ Estudio paleosismológico (metodologíay resultados)
- ✓ Conclusiones

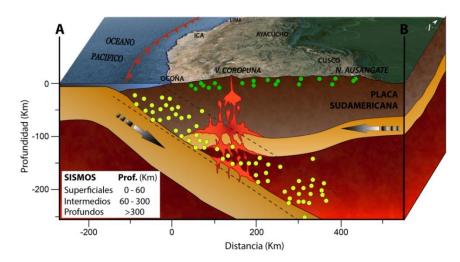


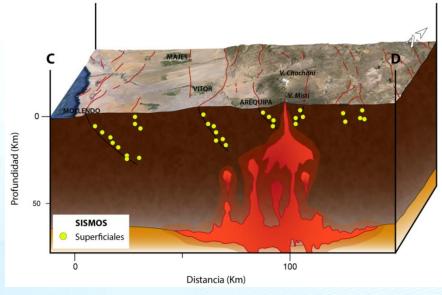




Fuente: IRIS Earthquake Browser



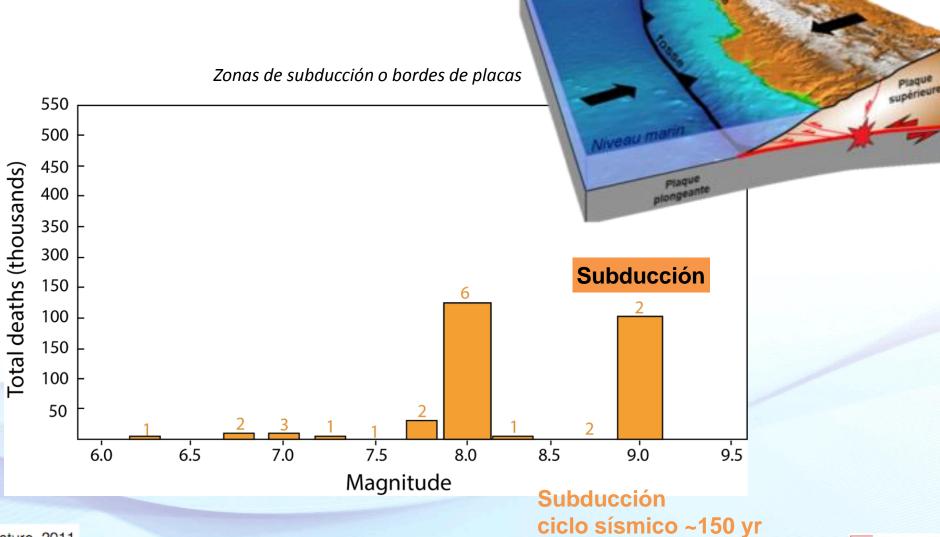








Peligro sísmico: Subducción versus Fallas activas



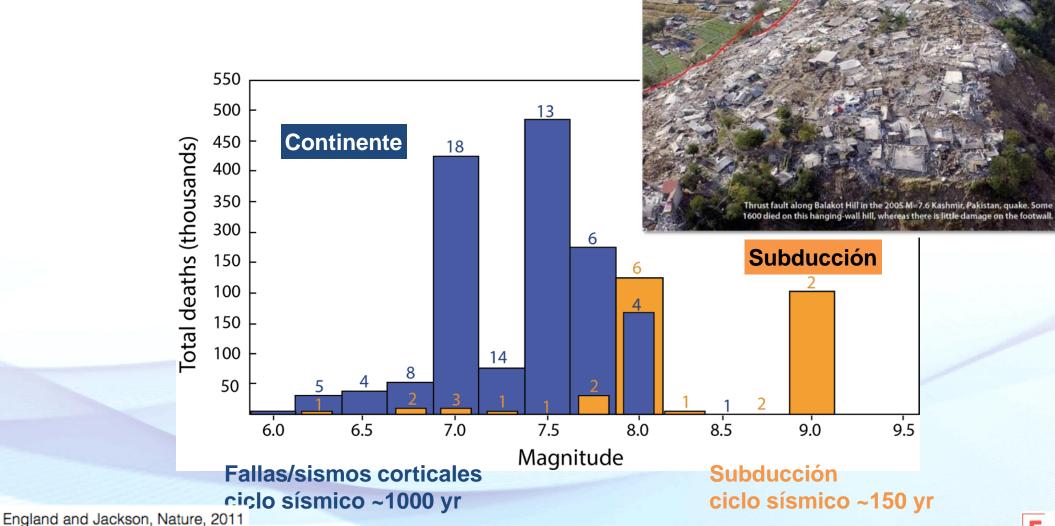
England and Jackson, Nature, 2011

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Bulletin of the Seismological Society of America

Vol. 44

APRIL, 1954

No. 2A

THE CUSCO, PERU, EARTHQUAKE OF MAY 21, 1950*

By George E. Ericksen, Jaime Fernández Concha, and Enrique Silgado

On May 21, 1950, an earthquake seriously damaged more than half the buildings of Cusco, Peru. It occurred without warning of foreshocks and was followed by only a few weak aftershocks. The pricentral area, within which buildings were extensively damaged, covered only about 12 square kilometers within the Cusco Basin, and minor damage was limited to a surrounding area of less than 500 square kilometers. The earthquake was felt over a total area of only 15,000 to 16,000 square kilometers and therefore may be considered local in effect. The depth of focus probably was between 8 and 9 kilometers.

The maximum intensity in the epicentral area was estimated to be VII on the Modified Mercalli scale, and the acceleration was less than 300 gals. The extensive damage in the epicentral area is attributed largely to poor construction and the age of buildings. The buildings that were most seriously damaged were colonial churches 250 to 350 years old, old adobe houses, and new houses of combined adobe and rock or brick construction. Many of the buildings that were destroyed in this earthquake had previously been damaged during an earthquake in 1941.

From a summary of the geology of the Cusco Basin it is concluded that the most extensive damage was to buildings resting on thick water-saturated alluvial gravels.

The earthquake was probably of tectonic origin, but there is no positive evidence of active faulting. However, an extensive fissure zone in unconsolidated sediments at the south side of the Cusco Basin may be related to a bedrock fault beneath these sediments.

Introduction

The Cusco earthquake of May 21, 1950, was felt over an area of 15,000 to 16,000 square kilometers between latitudes 13° 13′ and 14° 30′ S and longitudes 71° 25′ and 72° 55′ W. Cusco is in the nothern part of this area, at latitude 13° 31′ S and longitude 71° 59′ W, and is approximately 600 airline kilometers southeast of Lima, the capital of Peru. (See fig. 1.)

An official commission of two geologists and one seismologist was dispatched by the Instituto Nacional de Investigación y Fomento Mineros del Perú and the U. S. Geological Survey to study the seismic effects and damage in the Cusco area. From May 28 to June 4 the group made the reconnaissance studies on which this report is based.

Ing. Carlos Jiménez, chief of the Departmental Office of the Dirección de Caminos, Ministerio de Fomento del Perú, supplied the writers with cars for reconnaissance trips in the Cusco area; he and Ing. Humberto Guerra, Sub-Director de Arquitectura, gave much information that greatly aided the field study. Ing. Luís Dorich, Director of the Oficina Nacional de Plancamiento y Urbanismo del Perú,

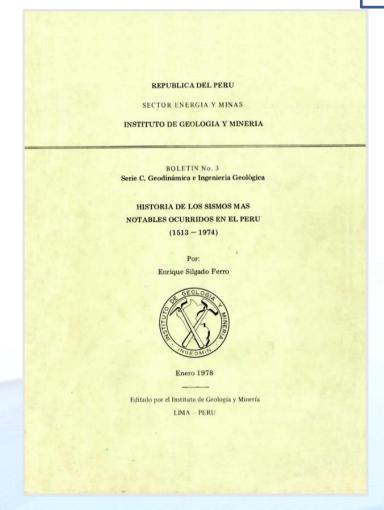
* Manuscript received for publication January 2, 1952.

1 Cuzco, in most English-language reference works.

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[97]

Ericksen et al. (1954)

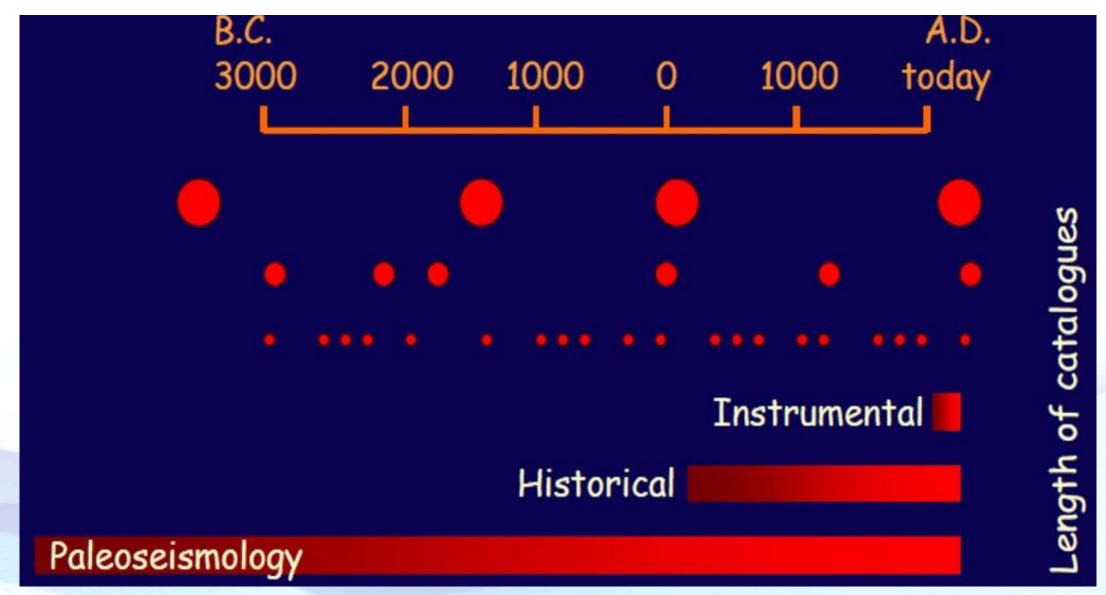


Silgado (1978)

Mapa neotectónico del Perú Océano Pacífico Fuente: Geocatmin - INGEMMET

EL PERÚ PRIMERO

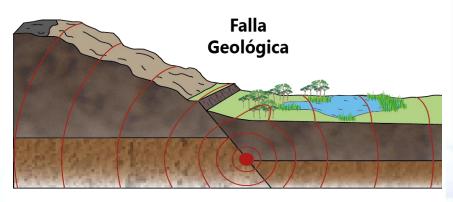












Fuente sismogénica??

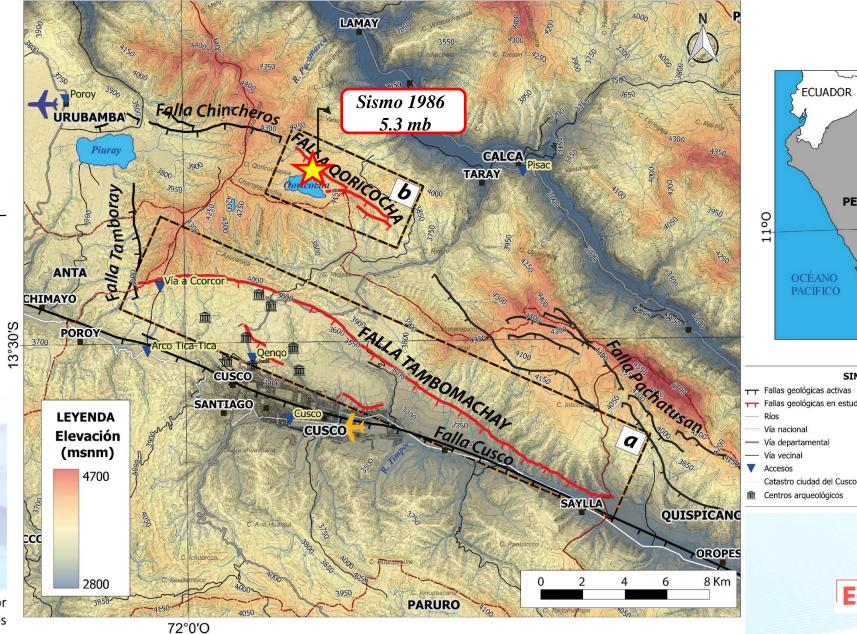
Silgado (1978)

EL PERÚ PRIMERO



Ubicación

Latitud (°)	Longitud (°)
13°26′44″S	72°0′36″O
13°29′5″S	72°1′42″O
13°34′39″S	71°48′48″O
13°32′27″S	71°47′47″O





/ía vecinal



Poblados importantes

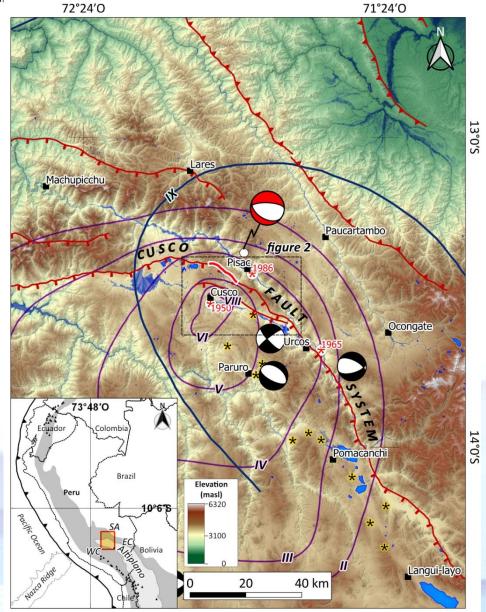
Límite provincial

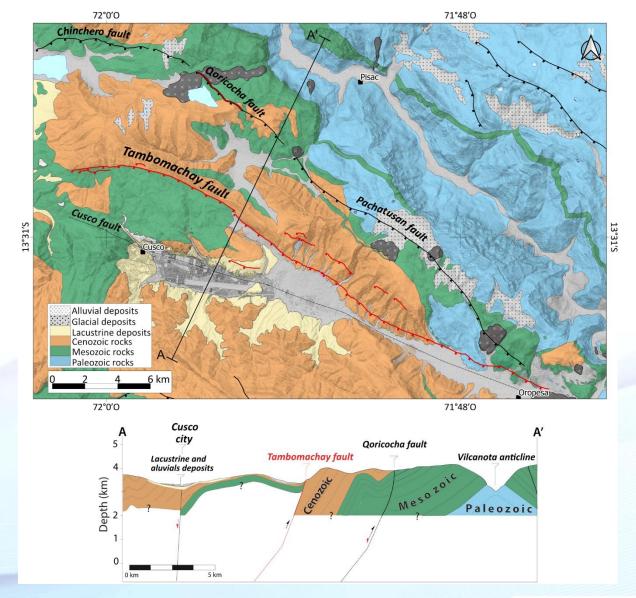
Zona de estudio

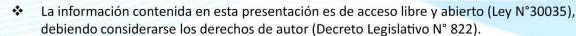
Lagos

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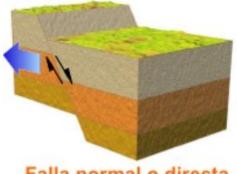






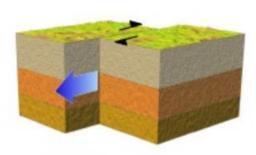
ESTUDIO GEOMORFOLÓGICO

1) Conocer la geometría y cinemática de la falla

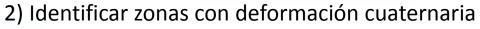


Falla normal o directa

Falla inversa



Falla de desgarre





3) Identificar zonas potenciales para el estudio paleosismológico





Data topográfica de alta resolución

Ortomosaico 1956 (2.5 m/px)

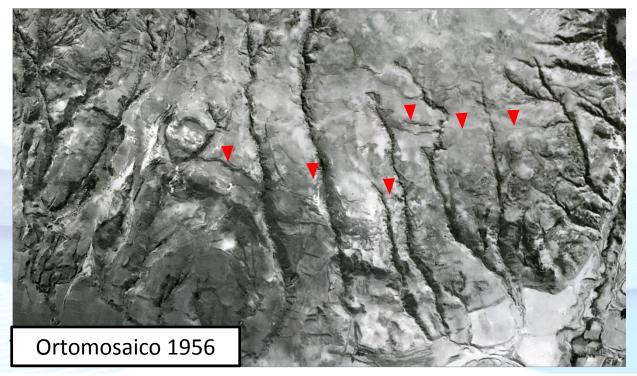












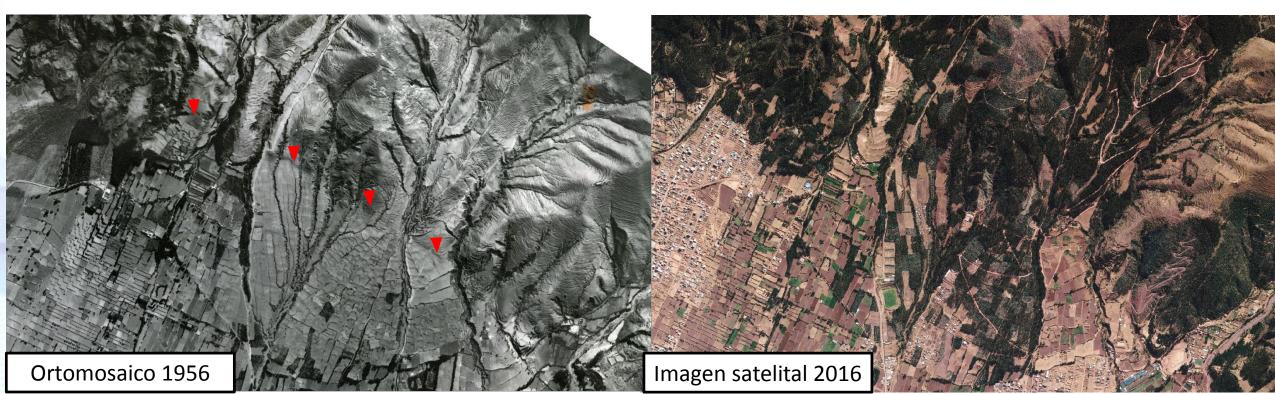






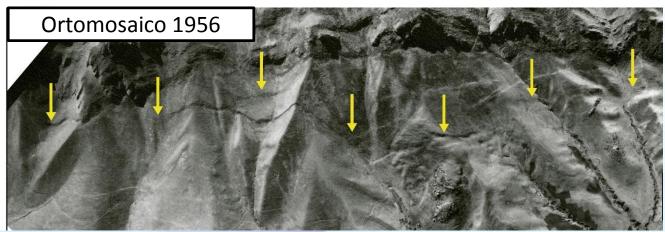


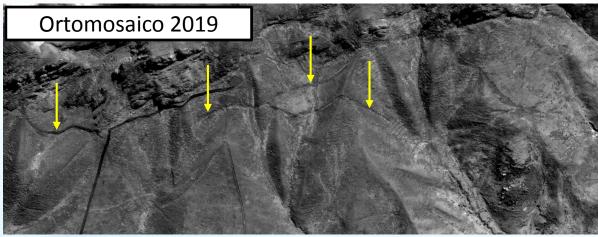






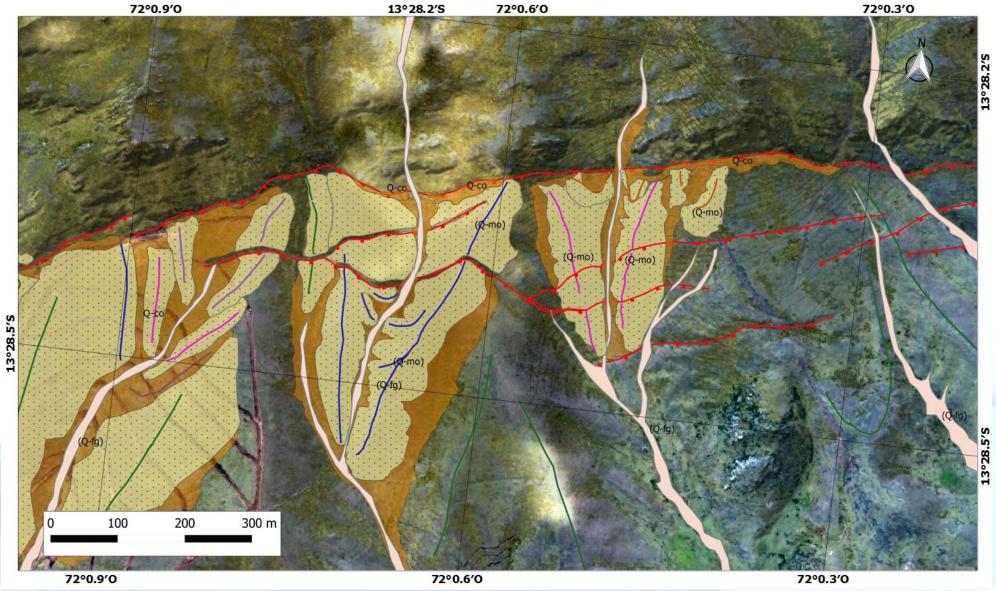


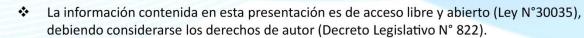






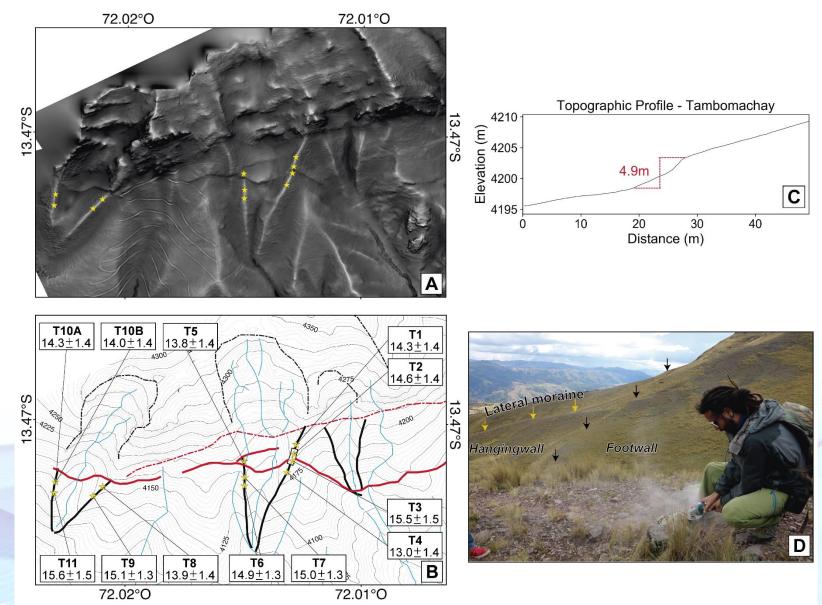














debido a estas observaciones realizamos el levantamiento de una columna estratigráfica.

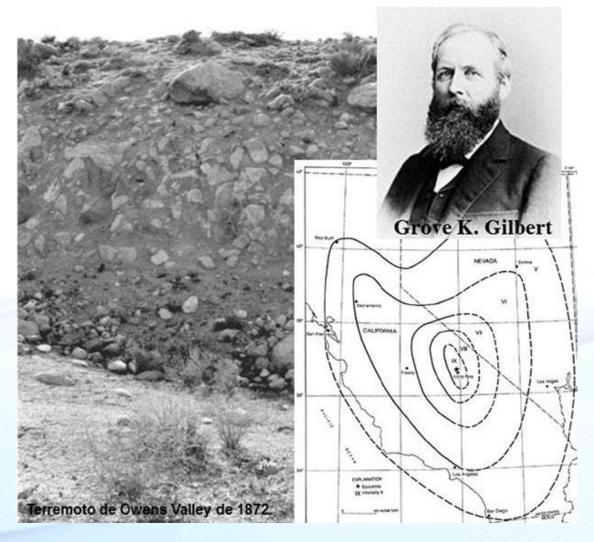


¿Qué es la paleosismología?

Herramienta geológica que estudia los terremotos ocurridos en el pasado.

(Yeats, 1996)

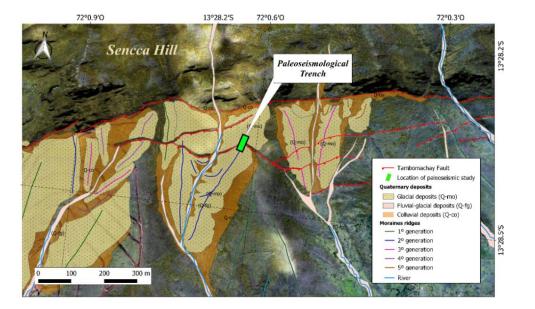






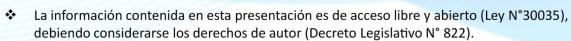






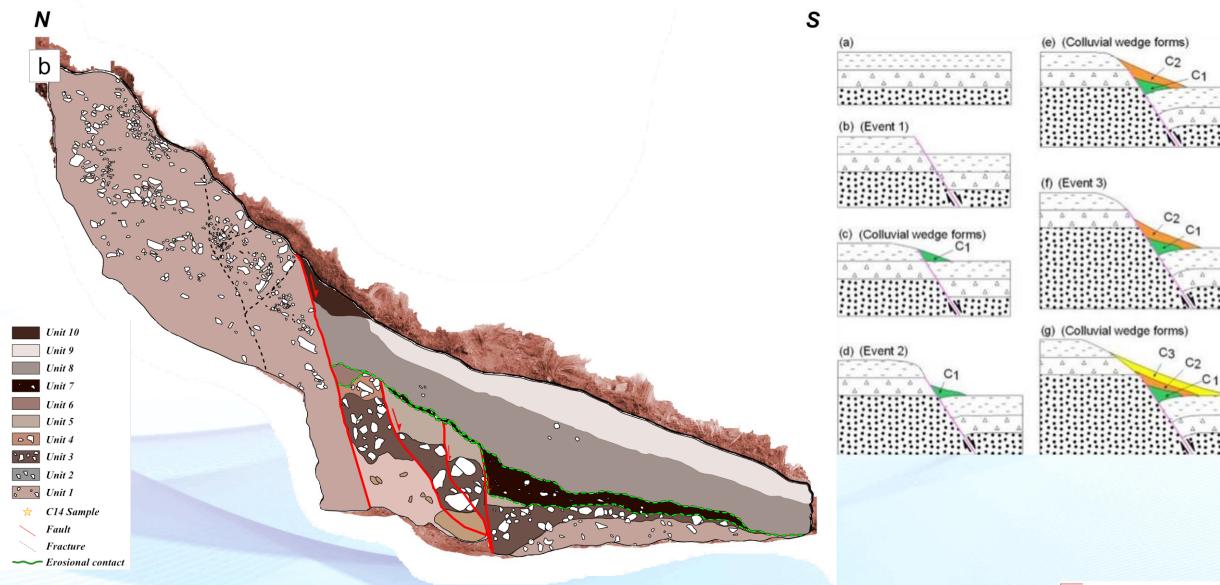


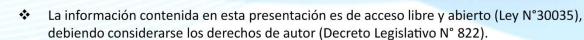






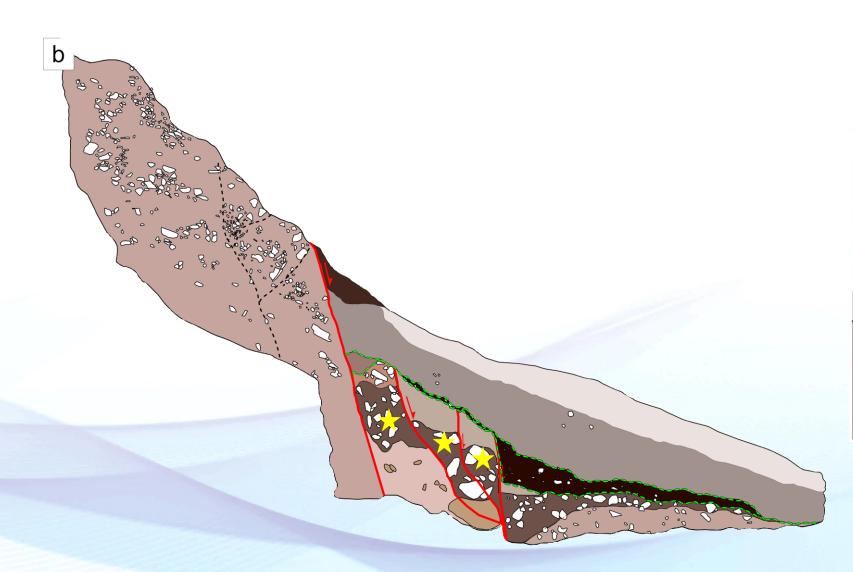


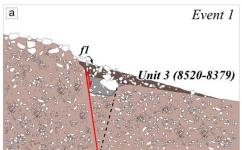


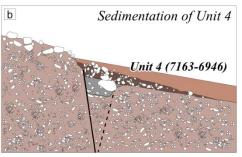


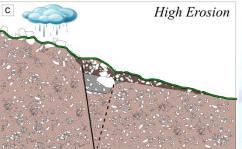






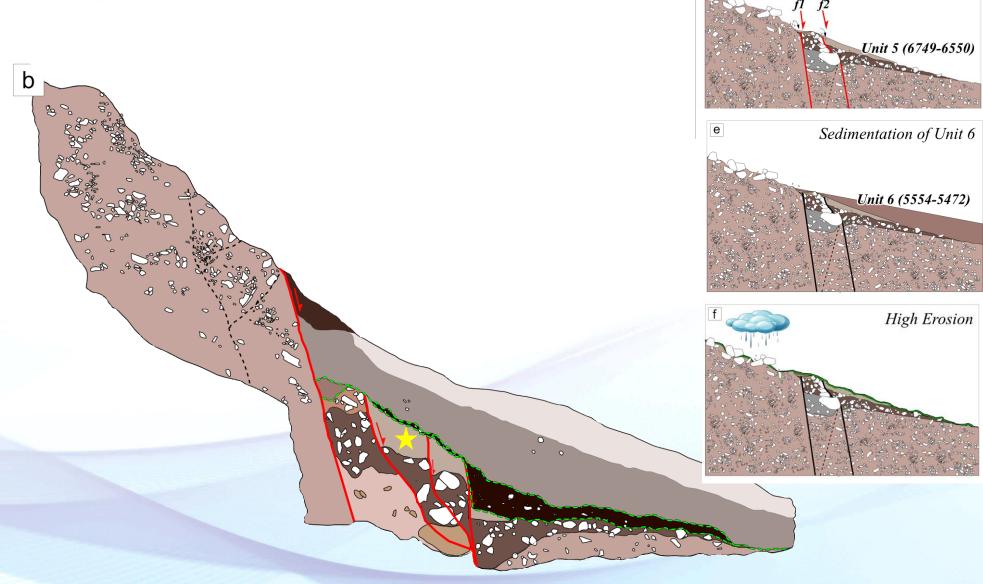








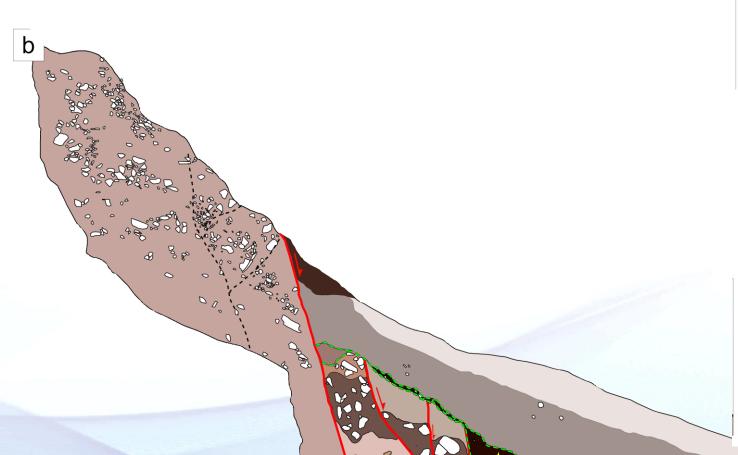


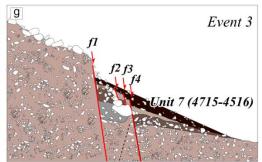


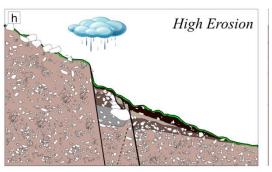


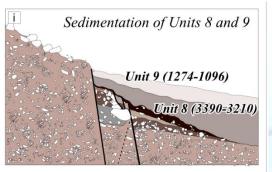
Event 2



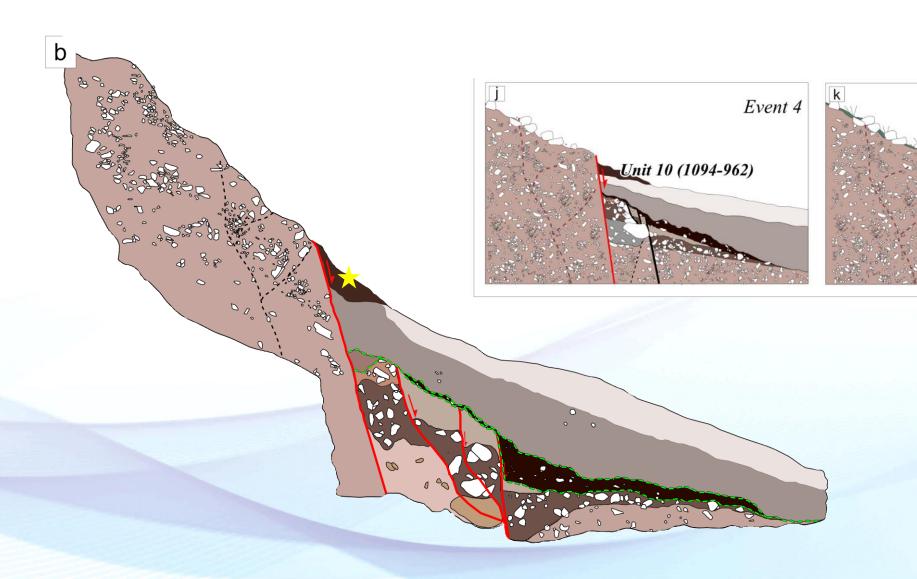














Current view



Cálculo de Paleomagnitudes

	Slip Number of		Coefficients and Standard Errors		Standard	Correlation	Magnitude	Displacement
Equation*		Events		b (sb)	Deviation s	Coefficient r	Range	Range (km)
$\mathbf{M} = a + b * \log (MD)$	SS	43	6.81(0.05)	0.78(0.06)	0.29	0.90	5.6 to 8.1	0.01 to 14.6
	$\{R\ddagger$	21	6.52(0.11)	0.44(0.26)	0.52	0.36	5.4 to 7.4	0.11 to 6.5}
	N	16	6.61(0.09)	0.71(0.15)	0.34	0.80	5.2 to 7.3	0.06 to 6.1
	All	80	6.69(0.04)	0.74(0.07)	0.40	0.78	5.2 to 8.1	0.01 to 14.6
$\log (MD) = a + b * M$	SS	43	-7.03(0.55)	1.03(0.08)	0.34	0.90	5.6 to 8.1	0.01 to 14.6
	{ <i>R</i>	21	-1.84(1.14)	0.29(0.17)	0.42	0.36	5.4 to 7.4	0.11 to 6.5}
	N	16	-5.90(1.18)	0.89(0.18)	0.38	0.80	5.2 to 7.3	0.06 to 6.1
	All	80	-5.46(0.51)	0.82(0.08)	0.42	0.78	5.2 to 8.1	0.01 to 14.6
$\mathbf{M} = a + b * \log (AD)$	SS	29	7.04(0.05)	0.89(0.09)	0.28	0.89	5.6 to 8.1	0.05 to 8.0
	$\{R$	15	6.64(0.16)	0.13(0.36)	0.50	0.10	5 8 to 7 4	0.06 to 1.5}
	N	12	6.78(0.12)	0.65(0.25)	0.33	0.64	6.0 to 7.3	0.08 to 2.1
	All	56	6.93(0.05)	0.82(0.10)	0.39	0.75	5.6 to 8.1	0.05 to 8.0
$\log (AD) = a + b * M$	SS	29	-6.32(0.61)	0.90(0.09)	0.28	0.89	5.6 to 8.1	0.05 to 8.0
	$\{R$	15	-0.74(1.40)	0.08(0.21)	0.38	0.10	5.8 to 7.4	0.06 to 1.5}
	N	12	-4.45(1.59)	0.63(0.24)	0.33	0.64	6.0 to 7.3	0.08 to 2.1
	All	56	-4.80(0.57)	0.69(0.08)	0.36	0.75	5.6 to 8.1	0.05 to 8.0

E	vento	AD (m)	Mw
	1	0.430	6.542
	2	0.763	6.704
	3	1.314	6.857
	4	0.430	6.542

Tomado de Wells & Coppersmith (1994)





CONCLUSIONES

- ✓ El trabajo demuestra que el análisis de imágenes satelitales y fotografías aéreas de alta resolución nos permiten caracterizar la evidencia geomorfológica de fallas activas; tales como la geometría y cinemática de la falla, identificar zonas con deformación reciente y potenciales para estudios paleosismológicos.
- ✓ Con el estudio paleosismológico logramos:
 - Afirmar que la falla muestra evidencias de deformación holocena.
 - Identificar 4 eventos sísmicos ocurridos en los últimos 8520 años.
 - Determinar los desplazamientos por evento, el que varía de entre 0.43 a 1.31 m.
 - Determinar que la magnitud de los terremotos generados es mayor a los 6.7 Mw.
 - Estimar un período de recurrencia de ~2000 años.
 - Por primera vez calculamos tasas de desplazamiento para esta falla, la cual varía entre 0.22 a 0.36 mm/año.
 - Determinar la edad del último evento, el cual aconteció en el año 900 d.C. o hace 1000 años.
- ✓ La integración de los resultados generados a partir del estudio geomorfológico junto a los datos obtenidos en la trinchera y la geocronología nos permiten hacer una caracterización completa de fuentes sismogénicas (fallas activas), información que posteriormente será ultilizada para la evaluación del peligro sísmico en una región.

